Vertical Distribution of Bio-Optical Properties of the Azov – Black Sea Basin Waters in April – May, 2019

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Purpose. The work is aimed at investigating spatial distribution of the chlorophyll *a* concentration and the spectral light absorption coefficients by all optically active components in the Azov and Black seas in spring, when the seawater hydrophysical structure changes.

Methods and Results. The data collected in the 106th scientific cruise of R/V Professor Vodyanitsky in April 19 - May 10, 2019 were used. The chlorophyll a concentration was measured by the spectrophotometric method. The spectral light absorption coefficients were determined in accordance with the NASA protocol 2018. The optical measurements were performed using the dualbeam spectrophotometer Lambda 35 (PerkinElmer). It was shown that in the surface layer of the Black Sea, the chlorophyll a concentration varied from 0.21 to 1.2 mg/m³. At some stations in the deep-water region, the increased values of this parameter were observed in the lower part of the euphotic zone that was associated with the beginning of seasonal water stratification due to the surface water heating. At these stations, the phytoplankton absorption spectra were more smoothed in the lower part of the euphotic zone than those in the upper layer. In the deep-water region, the non-algal particles contribution to the total particulate light absorption at wavelength 438 nm changed with depth from 40 ± 15 % at the surface to 29 ± 12 % near the bottom of the euphotic zone; whereas in the coastal waters this parameter was almost unchangeable within the water column (54 \pm 11 %). No significant change of the colored dissolved organic matter contribution to the total light absorption with depth was revealed (69% on average). In the Sea of Azov, vertical distribution both of the chlorophyll a concentration (6.2 mg/m³ on average) and the spectral light absorption coefficients by all the optically active components was uniform. The non-algal particles contribution to the particulate light absorption was 40 ± 14 %, and the colored dissolved organic matter contribution to the total light absorption constituted 52 ± 6 %.

Conclusions. New data on spatial distribution of the chlorophyll *a* concentration and the spectral light absorption coefficients by the optically active components in the Black and Azov seas were obtained for the spring period when the seawater hydrophysical characteristics changed.

Keywords: chlorophyll *a*, spectral light absorption coefficients, non-algal particles, colored dissolved organic matter, Black Sea, Sea of Azov.

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Introduction

active components of the The optically aquatic environment are phytoplankton, non-algal particles and colored dissolved organic matter. The distribution of phytoplankton in the water column is influenced by irradiance, temperature and nutrient availability. Analysis of the spatial distribution of chlorophyll a concentration and spectral light absorption by all optically active components revealed their variability regularities depending on the hydrophysical factors during the spring hydrological structure changes. These regularities are necessary for the development of regional model of chlorophyll *a* concentration in the surface layer [1], vertical distributions of chlorophyll *a* concentration [2] and spectral model of downwelling irradiance [3] in the Black and Azov seas, taking into account the seasonal characteristics of the environment.

The work is aimed to study spatial distribution of the chlorophyll *a* concentration and the spectral light absorption coefficients by all the optically active components in the Azov and Black seas in spring, when the seawater hydrological structure changes.

Material and methods

The measurements of the spectral bio-optical properties of the Azov and Black sea basin waters were carried out in the 106th scientific cruise of R/V *Professor Vodyanitsky* in April 19 – May 10, 2019 (Fig.1).



F i g. 1. Map of location of bio-optical stations in the 106^{th} cruise of R/V "Professor Vodyanitsky", April 19 – May 10, 2019

Water samples (4 L) were collected at various depths within the 0–70 m layer at daytime stations with Niskin bottle. The sampling depths were chosen based on the vertical profiles of temperature, chlorophyll *a* fluorescence, penetrating photosynthetically available radiation (PAR), and water transparency. Water transparency was assessed by Secchi disk depth visibility (Z_s). Vertical profiles of temperature, chlorophyll *a* fluorescence and PAR were obtained using an Idronaut Ocean Seven 320Plus CTD. 526 PHYSICAL OCEANOGRAPHY VOL. 27 ISS. 5 (2020) The euphotic zone (Z_{eu}) in the Black Sea was determined as depth with 1% of PAR incident on the sea surface. In the Azov Sea, where the PAR sensor was not used, Z_{eu} was calculated based on the dependence of light (I) attenuation with depth (z):

$$I(z) = I(0) \exp(-k_{\rm d} z),$$
 (1)

where k_d is the diffuse light attenuation coefficient on the average for the Z_{eu} layer for the PAR. The euphotic zone depth was calculated based on the Equation (1):

$$Z_{\rm eu} = 4.6/k_{\rm d}$$
.

Values of k_d ware calculated based on the Z_s values, using the dependence [4, p. 68].

To determine the chlorophyll *a* concentration in sum with phaeopigment (*Chl-a*) and the spectra of light absorption by suspended matter and phytoplankton pigments, 1–2 L water samples were gently vacuum filtered (~ 0.2 atm) through glass fiber filters (Whatman GF/F). Filters with a sample for the *Chl-a* determination were wrapped in an aluminum foil. Filters with samples for particulate light absorption spectra determination were placed in plastic holders. The samples were stored in a liquid nitrogen under the necessary conditions [5] until analysis on a laboratory. To determine the spectra of light absorption by colored dissolved organic matter (CDOM), 0.2 L samples were gently vacuum filtered (~ 0.2 atm) through a nucleopore filter (Sartorius) with a pore size of 0.2 μ m.

Chl-a in the sea was determined spectrophotometrically [6].

The determination of the spectral light absorption coefficients by suspended and dissolved organic matter was carried out in accordance with the modern NASA protocol [7]. Particulate light absorption was determined by the filter pad technique ("wet filter technique") [8, 9]. Phytoplankton light absorption $(a_{ph}(\lambda))$ was calculated by the difference between total particulate matter absorption $(a_p(\lambda))$ and non-algal particles absorption $(a_{NAP}(\lambda))$ according to the method described in [10]. Values of $a_{ph}(\lambda)$ were calculated from measured optical densities after correction for the path length amplification factor (β -correction) applying the quadratic equation described in [11, p. 146].

Optical measurements were made over the spectral domain from 350 to 750 nm (acetone extracts and suspended matter on a filter) and over the spectral domain from 250 to 750 nm (CDOM) with a dual-beam spectrophotometer LAMBDA 35 (PerkinElmer) equipped with a Spectralon integrating sphere.

Spectral distribution of light absorption coefficients by non-algal particles $(a_{\text{NAP}}(\lambda))$ and *CDOM* $(a_{\text{CDOM}}(\lambda))$ were described by the exponential functions

$$a(\lambda) = a (438) \exp(-S (\lambda - 438)),$$

where S is the spectral slope, nm^{-1} , obtained as result of fitting from 400 to 700 nm for NAP and from 350 to 500 nm for CDOM.

Results and discussion

The water transparency and euphotic zone depth in the Black Sea during the study period differed significantly between the stations. In the deep-water area of the Black Sea, the Z_s values varied from 11 to 19 m and was 14 m on average. Z_{eu} varied from 28 to 44 m and reached an average of 35 m. The coastal zone of

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the sea was characterized by lower water transparency: Z_s ranged from 5 to 11 m and averaged 7.5 m, while Z_{eu} varied from 15 to 28 m. The upper mixed layer (UML) depth, due to the spring warming of surface waters and the start of seasonal stratification, varied between stations from 3 to 50 m.

An important indicator of productivity and water quality is the chlorophyll a concentration – the main photosynthetically active pigment of phytoplankton. In the Black Sea, during the study period, the *Chl-a* value in the surface layer varied from 0.21 to 1.2 mg/m³ and was 0.60 mg/m³ on average. Significant differences were observed in the type of the vertical distribution of *Chl-a* values. Profiles with a homogeneous distribution of this parameter and profiles with an increase in *Chl-a* near the bottom of euphotic zone by 1.5–2.5 times in comparison with the *Chl-a* values in the surface sea layer were noted (Fig. 2).



F i g. 2. Vertical distribution of the measured chlorophyll *a* concentration taken together with phaeopigment *Chl-a* (green circles) and the *Chl-a* data from the CTD *Idronaut Ocean Seven 320 Plus* (green line), temperature *T* (blue line) and photosynthetically available radiation (PAR) (red line) at some stations in the deep part of the Black Sea in the second half of April – the first half of May, 2019

Since spring is a transitional period when the hydrological/hydrophysical vertical structure of the waters changes, the vertical distribution of phytoplankton pigments was different at the stations. At some deep-water stations, an almost uniform vertical distribution of chlorophyll a, typical for winter, was still observed within the UML, bounded from below by the maximum temperature gradient layer (Fig. 2). At other stations, where, as a result of the start of seasonal waters stratification, the phytoplankton existence layer was divided by the thermocline into two quasi-isolated layers (UML and the layer below thermocline), a chlorophyll *a* maximum appeared below the thermocline. Phytoplankton in these layers of the euphotic zone existed under various environmental conditions. Phytoplankton in the UML was under conditions of higher PAR and temperature (Fig.2), but likely with less nutrient availability than in the layer below thermocline. These differences determine changes in the structural and functional characteristics of phytoplankton existing in the upper photosynthetic zone and in the lower part of the euphotic zone, separated by a seasonal thermocline. In particular, a decrease in irradiance with depth within the euphotic zone leads to an increase in the chlorophyll *a* content in the cells of planktonic algae inhabiting the deep layer below the thermocline [12, 13].

Thus, the beginning of the disturbance of the typical winter-uniform vertical Chl-a distribution and an increase in the chlorophyll a content in the lower part of the euphotic zone is associated with the beginning of the spring warming of surface waters and the process of the seasonal stratification formation, resulted in poor deep vertical mixing of waters.

In the spectra of light absorption by phytoplankton pigments $(a_{ph}(\lambda))$, two main peaks are noted – at wavelengths of 438 and 678 nm (Fig. 3). The ratio between the absorption coefficients in these peaks (*R*) in the surface layer of the Black Sea was 2.8 ± 0.4. In the lower part of the euphotic zone, the $a_{ph}(\lambda)$ spectra were smoother and the R ratio was 2.3 ± 0.3. Such a change in the light-absorbing capacity of phytoplankton may reflect a photoadaptive decrease in the proportion of accessory pigments carotenoids [14] with depth within the euphotic zone and a change in the size structure of phytoplankton community [15].

In the surface layer of the deep-water part of the Black Sea, the light absorption by non-algal particles at a wavelength of 438 nm (where the blue maximum of the $a_{ph}(\lambda)$ spectrum is located) averaged 0.017 ± 0.008 m⁻¹, which coincides with the data obtained for the winter period of the year [16]. In this case, the relative contribution of the non-algal particles $a_{NAP}(438)$ to the total light absorption by all suspended matter at a wavelength of 438 nm ($a_p(438)$) was, on average, 40 ± 15 %. In the layer from 20 to 60 m of the deep-water part of the Black Sea, the a_{NAP} (438) value was 0.014 ± 0.006 m⁻¹, the contribution of $a_{NAP}(438)$ to $a_p(438)$ was 29 ± 12 %. In the waters of the coastal part of the Black Sea, the $a_{NAP}(438)$ coefficient was two times higher (0.037 ± 0.019 m⁻¹) than in the deep-water part; the $a_{NAP}(438)$ quota in $a_p(438)$ was 54 ± 11 %. The same high contribution (~ 50%) of the $a_{NAP}(438)$ to the total absorption by suspended matter was previously obtained in coastal waters near Katsiveli [17].



F i g. 3. Spectra of the light absorption coefficients by total particles $(a_p(\lambda))$ (*a*), non-algal particles $(a_{NAP}(\lambda))$ (*b*) and phytoplankton pigments $(a_{ph}(\lambda))$ (*c*) on the Black Sea surface in the second half of April – the first half of May, 2019

The spectral distribution $a_{\text{NAP}}(\lambda)$ is described by an exponential function (Fig. 3). Analysis of the exponential coefficient (S_{NAP}) variability in the Black Sea area revealed its weak variability both between stations and with depth. The S_{NAP} value averaged 0.009 \pm 0.001 nm⁻¹, which is consistent with the data obtained earlier for other regions of the World Ocean [18].

The CDOM light absorption coefficient at a wavelength of 438 nm $(a_{CDOM}(438))$ in the surface layer of the deep-water area of the Black Sea in the spring was $0.098 \pm 0.026 \text{ m}^{-1}$ (Fig. 4), and in the surface layer of the coastal part $-0.13 \pm 0, 05 \text{ m}^{-1}$. The $a_{CDOM}(438)$ contribution to the total light absorption by suspended and dissolved matter was $69 \pm 8 \%$.

The average value of the spectral slope coefficient S_{CDOM} in the surface layer of the Black Sea during the study period for the 350–500 nm domain was 0.017 \pm 0.002 nm⁻¹. This does not contradict the reference data, giving S_{CDOM} values for different World Ocean areas in the range from 0.014 to 0.025 nm⁻¹ [19].

It is known that the shape (S_{CDOM}) of CDOM light absorption spectra is dependent on the chemical content of colored dissolved organic compounds: the ratio of high- and low-molecular fractions [20]. The CDOM composition in terms of the ratio of high- and low-molecular-weight organic compounds is characterized by the *Sr* value – the ratio of exponential coefficients for the wavelength domains of 275–295 and 350–400 nm [21]. It is shown that the *Sr* value in the surface layer of the deep-water and coastal parts of the Black Sea in spring was the same and averaged 1.7 ± 0.2 .



Fig. 4. Spectra of the light absorption coefficients by colored dissolved organic matter $(a_{CDOM}(\lambda))$ in the surface layer of the Black Sea deep part (a), in the surface layer of the Black Sea coastal part (b) and in the Azov Sea waters (c) in the second half of April – the first half of May, 2019



F i g. 5. Spectra of the light absorption coefficients by particles $(a_{p}(\lambda))$ (*a*), non-algal particles $(a_{NAP}(\lambda))$ (*b*) and phytoplankton pigments $(a_{ph}(\lambda))$ (*c*) on the surface and in the water column of the Sea of Azov in the second half of April – the first half of May, 2019 PHYSICAL OCEANOGRAPHY VOL. 27 ISS. 5 (2020) 531

The low water transparency was noted in the Azov Sea in spring. The depth Z_s was ~ 2.5 m. Accordingly, the depth of the euphotic layer was ~ 9 m at the depth of the sampling stations from 10 to 12 m. That is, euphotic layer was almost the entire water column down to the bottom. An almost uniform vertical distribution of chlorophyll *a* concentration was observed. The *Chl-a* value in the sea surface layer varied between stations from 4.9 to 8.8 mg/m³.

The $a_{ph}(\lambda)$ in Azov Sea was an order of magnitude larger than in the Black Sea (Fig. 3, 5). The R values in the Azov Sea surface layer and water column was less than in the Black Sea waters, and was 2.00 ± 0.26 on average (Fig. 3, 5).

The Azov Sea waters differed from the waters of the Black Sea by an order of magnitude larger values of a_{NAP} (438) (0.11 ± 0.04 m⁻¹). The relative contribution of a_{NAP} (438) to a_{p} (438) was 40 ± 14 %, on average, as in the surface layer of the deep-water part of the Black Sea. S_{NAP} ratio was the same as in the Black Sea.

The $a_{\text{CDOM}}(438)$ coefficient in the Azov Sea was almost three times higher than in the Black Sea and was $0.310 \pm 0.058 \text{ m}^{-1}$ on average (Fig. 4). The corresponding $a_{\text{CDOM}}(438)$ contribution to the total light absorption by all suspended and dissolved matter was slightly less than in the Black Sea and equaled 52 ± 6 %. At the same time, slightly higher S_{CDOM} values were obtained in the Azov Sea waters for the domain of 350–500 nm (0.019 \pm 0.001 nm⁻¹) than in the Black Sea waters. In both the Black and Azov seas, the S_{CDOM} values are in the range corresponding to the reference data [19]. The S_r coefficient in the more eutrophic Azov Sea waters, on the contrary, was lower and amounted to 1.4 ± 0.1 .

Conclusion

New data on spatial distribution of the chlorophyll *a* concentration and the spectral light absorption coefficients by the optically active components in the Black and Azov seas were obtained for the spring period when the seawater hydrophysical characteristics changed. The revealed regularities will improve the accuracy of the existing regional models for the Black Sea and develop appropriate models for the Azov Sea.

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Elena Yu. Skorokhod - data processing

Nataliya A. Moiseeva - data processing

Ekaterina A. Zemlianskaia - sampling, drawing

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